Statistica Sinica Preprint No: SS-2017-0403	
Title	Detection and replenishment of missing data in marked point processes
Manuscript ID	SS-2017-0403
URL	http://www.stat.sinica.edu.tw/statistica/
DOI	10.5705/ss.202017.0403
Complete List of Authors	Jiancang Zhuang Ting Wang and Koji Kiyosugi
Corresponding Author	Jiancang Zhuang
E-mail	zhuangjc@ism.ac.jp

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Detection and replenishment of missing data in marked point processes

Jiancang Zhuang[1)], *Ting Wang*[2)], *and Koji Kiyosugi*[3)]

[1)] *Institute of Statistical Mathematics*

10-3 Midori-cho, Tachikawa, Tokyo 190-8562, Japan

[2)] *Department of Mathematics and Statistics, University of Otago, Dunedin, New Zealand*

[3)] *Organization of Advanced Science and Technology, Kobe University*

1-1 Rokkodai-cho, Nada-ku, Kobe 657-8501, Japan

August 13, 2019

Abstract

Records of geophysical events, such as earthquakes and volcanic eruptions, are usually modeled as marked point processes. These records often suffer from missing data, resulting in underestimations of the corresponding hazards. We propose a computational approach for replenishing data missing from the records of temporal point processes with time-separable marks. The proposed method is based on the notion that if such a point process is completely observed, it can be transformed into a homogeneous Poisson process, approximately on the unit square $[0, 1]^2$, by a biscale empirical transformation (BEPIT). This approach includes three key steps: (1) transforming the process onto $[0, 1]^2$ using the BEPIT, and finding a time–mark range that likely contains missing events; (2) estimating a new empirical distribution function based on the data in the time–mark range in which the events are supposed to be completely observed; and (3) generating events in the missing region. We test this method on a synthetic data set, and apply it to records of the

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

volcanic eruptions of the Hakone Volcano in Japan and the aftershock sequence following the 2008 Wenchuan Mw7.9 earthquake in Southwest China. The results show that this algorithm provides a useful way to estimate missing data and to replenish incomplete records of marked point processes. In addition, the replenished data provide estimates of the hazard function that are more robust.

1 Introduction

Many geophysical processes, such as earthquakes and volcanic eruptions, occur at random times and/or locations, and, thus, are described naturally by point-process models (e.g., Vere-Jones, 1970; Zhuang et al., 2002; Wang and Bebbington, 2012, 2013). Point-process models and their related theories are also widely used in fields such as crime, disease, and fire (Diggle and Rowlingson, 1994; Schoenberg et al., 2007; Mohler et al., 2011). Furthermore, advancements in the technology used to record these natural and social phenomena are yielding significantly greater amounts of data. However, the degree of completeness of these records varies, and in many cases, small events are often missed in the early period of observation. For example, smaller aftershocks are less likely to be recorded than are larger aftershocks during the period immediately following a large earthquake (Ogata and Katsura, 1993; Omi et al., 2013). Other examples include missing data in volcanic eruption records (Kiyosugi et al., 2015) and in the field of communication in social networks (Zipkin et al., 2015). Missing data limit our efficient use of these records, often resulting in biased estimates. However, statistical tools for analyzing incomplete point-process data are not well developed.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Geophysicists have been searching for reliable methods of obtaining earthquake catalogs that are more complete. For example, waveform-based detection methods for small earthquakes within an aftershock sequence have been proposed (e.g., Enescu et al., 2007, 2009; Peng et al., 2007; Marsan and Enescu, 2012; Hainzl, 2016). However, even these methods cannot recover all missing aftershocks. An alternative is to switch to energy-based descriptions (Sawazaki and Enescu, 2014); that is, rather than viewing earthquake occurrences as a process of events with different magnitudes, the process is regarded as a stream of energies released by earthquakes. However, methods related to such descriptions remain underdeveloped.

Based on the empirical law that the distribution of earthquake magnitudes follows the Gutenberg–Richter magnitude–frequency relation (Gutenberg and Richter, 1944), Ogata and others investigated why events were missing from earthquake catalogs (Ogata and Vere-Jones, 2003; Iwata, 2008, 2013, 2014). They used a Bayesian method to make probabilistic earthquake forecasts, with missing earthquakes taken into account (Ogata, 2006; Omi et al., 2013, 2014, 2015).

In most of the aforementioned studies, when dealing with missing events in a point process, the full structure of the model, or at least the distribution of marks, is assumed to be known. However, owing to incomplete records and other reasons, on most occasions, the information available on the process or the mark distribution is limited. Thus, a preferable method for evaluating the missingness should be based on as few assumptions as possible, especially when the temporal structure and the distribution of marks are unknown. Zhuang et al. (2017) used a stochastic algorithm to restore missing aftershocks in the aftershock sequences following several earthquakes in

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Kumamoto, Japan (April 14, 2016, $M6.5$; April 15, 2016, $M6.4$; April 16, 2016, $M7.3$). This method can be used to restore missing data in the records of a more general temporal point process with time-separable marks, using information from the parts of the process that are completely observed. In Zhuang et al. (2017), the mathematical background is not well addressed. In this study, we explain in detail the mathematics related to this fast algorithm and discuss its asymptotic properties.

In the following sections, we first introduce the biscale empirical probability integral transformation (BEPIT), and then analyze the completely observed process with time-separable marks after the transformation. Based on the results of this transformation, we restore the empirical distributions from an incomplete record using an iterative algorithm. We explain the algorithm using a simulated data set. Finally, we apply the algorithm to investigate the incomplete eruption record of the Hakone volcano in Japan, and the aftershock sequence of the Wenchuan Mw7.9 earthquake that occurred in Southwest China on May 28, 2008. The proofs of the consistency and asymptotic normality of the algorithm are given in Supplementary Material.

2 Concepts, methodology, and illustration

2.1 Mark-separable temporal point process and BEPIT

Mathematically, a marked temporal point process N is a random subset of discrete points on the space $\mathbb{R} \times \mathbb{M}$, say $\{(t_i, m_i) : i = 1, 2, \cdots, n\}$, which includes a finite or countable number of elements, and satisfies the following two conditions (Karr, 1991): (a) for any bounded subset $A \subset \mathbb{R}$, $\Pr\{N(A \times \mathbb{M}) \equiv \#[N \cap (A \times \mathbb{M})] \ < \infty\} = 1$, where $\#[\ ]$ represents the number of elements in a set; and (b) for each i, m_i is a random variable on $\mathbb{M}$. In

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

our study, we assume the following: (a) the marks are continuous random variables, and (b) the point process is simple (i.e., $\Pr\{\max_{t\in\mathbb{R}} N(\{t\}\times\mathbb{M}) \leq 1\} = 1$), such that there are no overlapping events on the time axis.

A marked temporal point process is often specified by its conditional intensity function, defined by

$$\lambda(t,m)\,\mathrm{d}t\,\mathrm{d}m = \mathbf{E}\left[N([t,t+\mathrm{d}t)\times(m,m+\mathrm{d}m) \mid \mathcal{H}_t\right], \tag{1}$$

where $\mathcal{H}_t$ denotes the history of N up to time t, but not including t. The conditional intensity can be decomposed as

$$\lambda(t,m) = \lambda_g(t)\, g(m|t),$$

where $\lambda_g(t) = \int_{\mathbb{M}} \lambda(t,m)\,\mathrm{d}m$ is called the conditional intensity of the ground point process N_g induced by N on $\mathbb{R}$, defined by $N_g(A) = N(A\times\mathbb{M})$, and $g(m|t)$ is the probability density function of the event mark at time t. An important property of the conditional intensity is that if a temporal point process N has conditional intensity $\lambda(t)$, then the transformation

$$t_i \to \tau_i = \int_0^{t_i} \lambda(u)\,\mathrm{d}u \tag{2}$$

transforms N into a Poisson process $N' = \{\tau_i : i = 1, 2, \cdots\}$ (see, e.g., Ogata, 1988; Schoenberg, 2003; Daley and Vere-Jones, 2003).

For the above conditional intensity, when the mark distribution is separable from the occurrence times, that is,

$$\lambda(t,m) = \lambda_g(t)\, g(m), \tag{3}$$

the marks of this point process are said to be time separable. Point-process models with time-separable marks are widely used in many research areas. In

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

seismology, most practical versions of earthquake forecasting models explicitly assume that the magnitude distribution is separable from time (see, e.g., Ogata and Zhuang, 2006; Zhuang et al., 2002, 2004; Zhuang, 2011; Werner et al., 2011; Ogata et al., 2013). In volcanology, Bebbington (2014) suggested that there is not enough evidence of a universal dependence of eruption size on time. In forecasting, time-independent size distributions are used frequently (e.g., Passarelli et al., 2010).

Other ways to specify point-process models include moment intensity functions, Papangelou intensities, and Palm intensities. Traditionally, when a point process is specified in one of these ways, it refers to a spatial point process. A point process can be completely determined by its likelihood (terminologically, the local Janossy density; see Daley and Vere-Jones, 2003, 2008). This gives the joint probability density/mass function of the total number and each location of the particles in the process, assuming that the particles are indistinguishable. The likelihood is also known (i.e., the point process is completely determined) if one of the following three is known: (1) the moment intensities of all orders, (2) the conditional intensity, and (3) the Papangelou intensity. Here, we refer to Daley and Vere-Jones (2003, 2008) and Møller and Waagepetersen (2003) for the relations between the Janossy density and three other types of intensities. In this study, the method used to replenish missing data in a marked point process does not depend on any specific form of conditional intensity. Therefore, it can be applied to spatial point processes as well if the ground space is one dimensional and the conditional intensity is mark separable.

Before testing for missing data in a record of a marked point process and replenishing the record, we need to know what a complete record looks like.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Given a series of independent and identically distributed (i.i.d.) observations on X, $x_1, x_2, \cdots, x_n$, for a fixed x, the empirical cumulative distribution function (cdf)

$$\tilde{F}_X(x) = \frac{1}{n}\sum_{i=1}^{n} \mathbf{1}(x_i < x)$$

converges almost surely to $F_X(x)$ and, thus, $\tilde{F}_X(X_j)$, for $j = 1, 2, \cdots, n$, converges to a unit uniform distribution. We call transformation $x \to \tilde{F}_X(x)$ the empirical probability integral transformation induced by $\{x_1, x_2, \cdots, x_n\}$. In a general marked point process N in $[0, T]$, the occurrence times of an arbitrary event may depend on the occurrence times and/or marks of other events. However, the empirical probability integral transformation still results in an approximate unit uniform distribution, because the transformation does not require an explicit formulation of the conditional intensity.

Suppose $N = \{(t_i, m_i) : i = 1, 2, \cdots, n\}$ is a realization of a temporal marked point process in a time–mark domain $[0, T] \times \mathbb{M}$, where $\mathbb{M}$ is the space of marks. Consider the following BEPIT:

$$\begin{array}{rccc} \Gamma_N : & [0,T]\times\mathbb{M} & \to & [0,1]\times[0,1] \\ & (t,m) & \to & (t',m') = \left(\tilde{F}(t), \tilde{G}(m)\right), \end{array} \tag{4}$$

where $\tilde{F}$ and $\tilde{G}$ are the empirical cdfs of $\{t_i : i = 1, 2, \cdots, n\}$ and $\{m_i : i = 1, 2, \cdots, n\}$, respectively. If the marks of the events in the process are separable from the occurrence times, then $\{t_i' : i = 1, 2, \cdots, n\}$ and $\{m_i' : i = 1, 2, \cdots, n\}$, which are the images of $\{t_i : i = 1, 2, \cdots, n\}$ and $\{m_i : i = 1, 2, \cdots, n\}$, respectively, form an approximately homogeneous Poisson process on $[0, 1] \times [0, 1]$. It is straightforward to show the independence between $\tilde{F}(t)$ and $\tilde{G}(m)$. Thus, given the total number of events N, the number of events in a cell of area $s \subseteq [0, 1] \times [0, 1]$ is a random variable from

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

a binomial distribution $B(N, s)$, which can be approximated by a Poisson distribution with mean Ns. The smaller s gets, the better this approximation becomes.

In the following discussion, we consider only mark-separable Poisson processes. This is because we can transform a more general process, say N, with a conditional intensity $\lambda(t, m)$, into a Poisson process N' with a constant intensity using the marked version of the transformation in (2), $(t_i, m_i) \in N \to (\tau_i, m_i) \in N'$, where $\tau_i = \int_0^{t_i} \int_{\mathbb{M}} \lambda(t, m)\, \mathrm{d}m\, \mathrm{d}t$. Because such a transformation does not change the chronological order of the events or the mark-separable property of the process, the BEPIT transforms N and N' into the same point patterns.

Example 1. *In Figure 1(a), we simulate a Poisson process N (the combination of dots and crosses) with a temporal rate $\lambda = 1$ on [0, 2,000], and marks that follow an exponential distribution with mean one; that is,*

$$g(x) = \begin{cases} \mathrm{e}^{-x}, & x > 0; \\ 0, & \textit{otherwise.} \end{cases}$$

Figure 1(b) shows that under transformation (4), N is transformed into an approximately homogeneous Poisson process, say N', which has rate $\lambda = 2,000$ and i.i.d. marks uniformly distributed in $[0, 1]$.

2.2 Detection of missing data

When events in part of an observed time–mark range are missing, deterministically or in probability, the separability between the occurrence times and the marks of the observed events is usually destroyed. In addition, the image of the observed N_{obs} mapped by the above BEPIT $\Gamma_{N_{\text{obs}}}$, as defined in (4), may not be a homogeneous process.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

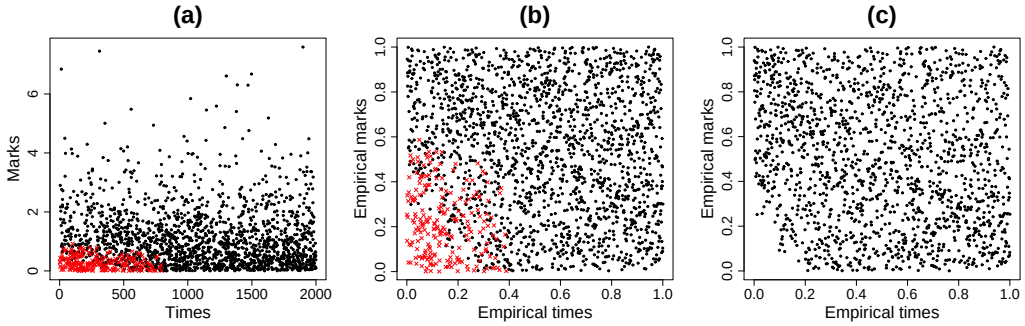

Figure 1: A synthetic data set of a marked point process. (a) Marks versus occurrence times. (b) Empirical marks versus empirical occurrence times of all synthetic events under the transformation Γ_N. (c) Empirical marks versus empirical occurrence times for the observed incomplete record under the transformation $\Gamma_{N_{\rm obs}}$. The crosses in (a) and (b) represent the missing events.

Example 2. *Consider the simulated data in Example 1 (Figures 1(a)). Assume the missing probability is*

$$\begin{aligned} q(t,m) &= \Pr\{\textit{an event occurring at } (t,m) \textit{ is missing}\} \\ &= \begin{cases} \min\left[1, \frac{(1000-t)(1-m)}{800}\right], & \textit{if } 0<t<800,\ m<0.3, \\ 0, & \textit{otherwise.} \end{cases} \end{aligned} \tag{5}$$

If we thin the original process N (the combination of the crosses and dots) in Figure 1(a) with this missing probability, then the crosses are deleted (i.e., they are missing from the record). Denote the remaining events (i.e., the observed process) as $N_{\rm obs}$. Figure 1(c) shows that the image of the observed data of the process under the BEPIT $\Gamma_{N_{\rm obs}}$ is not homogeneous.

In the above bi-scale transformation, we do not need to know the exact forms of $g(m)$, λ_g, or q. This method relies only on the conditions that the original process is mark separable, and that the process of missing events is time- and mark-dependent. Thus, for a temporal point process N with time-separable marks, we can test whether data are missing from its observed record, $N_{\rm obs}$, by testing the homogeneity of the image $\Gamma_{N_{\rm obs}}(N_{\rm obs})$ of the

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

observed data N_{obs} in the bi-scale transformed domain, when the missing values are time- and mark-dependent. After using the BEPIT $\Gamma_{N_{\text{obs}}}$ to map N_{obs} onto $[0,1]^2$, we divide the overall area of $[0,1]^2$ into L sub-regions of equal areas, that is, $L = L_1 \times L_2$ cells. Here, L_1 is the number of cells along the transformed time domain, and L_2 is the number of cells along the transformed mark domain. Then, we calculate the following statistics:

$$R = \frac{\min\{C_1, C_2, \cdots, C_L\}}{\max\{C_1, C_2, \cdots, C_L\}}, \text{ and } D = \max\{C_1, C_2, \cdots, C_L\} - \min\{C_1, C_2, \cdots, C_L\}, \tag{6}$$

where $C_1, C_2, \cdots, C_L$ are the numbers of events falling within each of the L cells. These two statistics are analogous to the test statistics for homogeneous multinomial distributions, where "homogeneous" means that each category of the possible outputs has the same probability (Johnson, 1960; Johnson and Young, 1960; Corrado, 2011).

Suppose that $[0,1]^2$ is divided into $L = L_1 \times L_2$ cells with equal areas; that is, $[0,1]^2 = \bigcup_{j=1}^{L_2} \bigcup_{i=1}^{L_1} [(i-1)/L_1, i/L_2] \times [(j-1)/L_2, j/L_2)$, where L_1 and L_2 are positive integers. For any point process N on $[0,1]^2$, if N is a homogenous Poisson process, then the numbers of events in the above L cells, $C_1, C_2, \cdots, C_L$, form a homogeneous $(n, \mathbf{p})$-multinomial random vector, with $\mathbf{p} = (1/L, 1/L, \cdots, 1/L)$. However, if N is obtained by applying the BEPIT to a completely observed mark-separable point process, then the row sum of C_i in the kth row ($1 \le k \le L_1$) and the column sum of C_i in the jth column ($1 \le j \le L_2$) are fixed to $\lfloor kn/L_1 \rfloor - \lfloor (k-1)n/L_1 \rfloor$ and $\lfloor jn/L_2 \rfloor - \lfloor (j-1)n/L_2 \rfloor$, respectively, where $\lfloor x \rfloor$ denotes the integer part of x, and n is the total number of events in N. Such constraints do not hold for the homogeneous multinomial distribution. Because the distributions of R

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

and D are complicated, we obtain them by simulation, as follows: (1) with n fixed, simulate n events uniformly distributed in $[0,1]^2$; (2) apply the BEPIT to these n simulated events; (3) with the specified parameters, L_1 and L_2, calculate R and/or D for the transformed points.

Example 3. *We use a simulation to test for missing data in the original and the thinned point processes, as shown in Figures 1(a) and (c), respectively. We simulate 500,000 sequences of the marked Poisson process defined in Example 1, with the number of events in each simulation the same as those in Figure 1(a). For each simulated sequence, we apply the BEPIT in (4), which results in an image similar to the combination of the crosses and the dots in Figure 1(b). Then, we divide the unit square image into five-by-five cells with equal sizes, and calculate R and D, as defined in (6). Next, we plot the empirical cdf of the 500,000 values of R and D, as shown in Figure 2(a). To test the thinned process, we simulate further 500,000 sequences of the marked point process, with the total number of events in each simulation the same as those in Figure 1(c). The cumulative distributions of R and D are shown in 2(b). We can see that the hypothesis of no missing data in the observed (thinned) process can be rejected, with a significance level below 0.001 ($p \leq 2 \times 10^{-6}$, Figure 2(b)). Meanwhile, for the original process,* the p-values associated with R and D (0.396 and 0.700, respectively) provide no evidence for rejection.

2.3 Imputation method and algorithm

We start with a heuristic example to explain the algorithm. As shown in Figure 3, suppose that N is a homogeneous point process on $[0,1]^2$, and that events in the domain S are completely unobservable. Let $N_{\text{obs}} = \{(x_i, y_i) :$

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

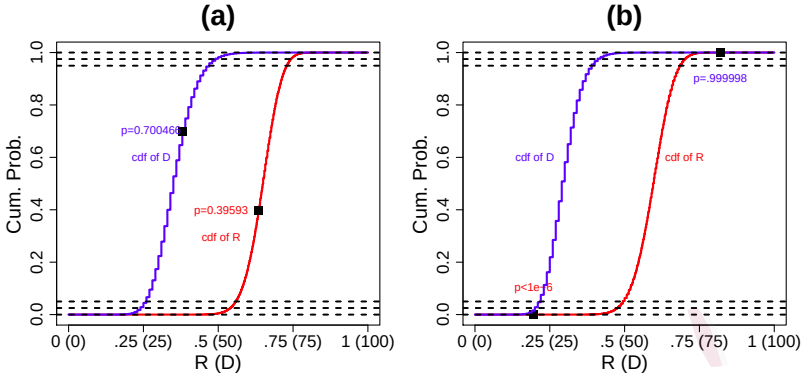

Figure 2: Statistical tests of the existence of missing data on (a) all events and (b) the observed events in the synthetic point process, with cdfs of R and D. R and D are defined in (6), with $L = L_1 \times L_2$, $L_1 = L_2 = 5$. The cdfs in (a) and (b) are obtained from 500,000 simulations with the same numbers of events as in Figures 1(a) and (c), respectively. The black dots in (a) and (b) are the statistics R and D, calculated for the original process in 1(a) and (c), respectively.

$(x_i, y_i) \in N \backslash S\}$. Then, the empirical distributions of the x- and y-coordinates are, respectively,

$$\tilde{F}_X(x) = \frac{\sum_{i:(x_i,y_i)\in N\backslash S} w_{x,i} I(x_i \le x)}{\sum_{i:(x_i,y_i)\in N\backslash S} w_{x,i}} \tag{7}$$

and

$$\tilde{F}_Y(y) = \frac{\sum_{i:(x_i,y_i)\in N\backslash S} w_{y,i} I(y_i \le y)}{\sum_{i:(x_i,y_i)\in N\backslash S} w_{y,i}}, \tag{8}$$

where

$$w_{x,i} = \frac{1}{1 - \int_0^1 I((x_i, y) \in S)\mathrm{d}y}, \qquad w_{y,i} = \frac{1}{1 - \int_0^1 I((x, y_i) \in S)\mathrm{d}x}. \tag{9}$$

In most cases, N is not homogeneous in $[0,1]^2$, and the variation of the event density in S should be considered. Equation (9) should then be

$$w_{x,i} = \frac{1}{1 - \int_0^1 I((x_i, y) \in S)\mathrm{d}F_y(y)}, \qquad w_{y,i} = \frac{1}{1 - \int_0^1 I((x, y_i) \in S)\mathrm{d}F_X(x)}. \tag{10}$$

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

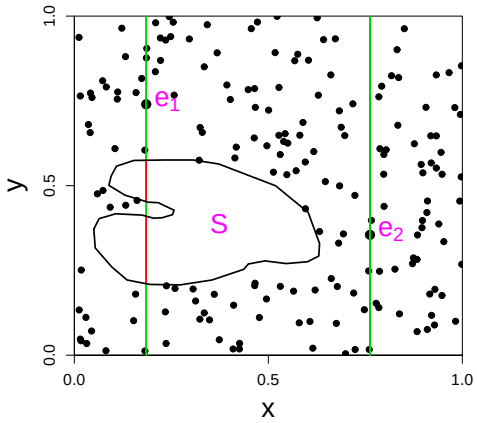

Figure 3: A heuristic estimation of the empirical distribution with missing points. Suppose that, among events $e_i = (x_i, y_i)$, for $i = 1, 2, \cdots, N$, events that fall in S cannot be observed. To estimate the empirical distribution $\tilde{F}_X(x)$ of x_i, for $i = 1, 2, \cdots, N$, weights need to be assigned to each observed point. That is, when N is uniform, $\tilde{F}_X(x) = \sum_{i=1}^{N} w_{x,i} I(x_i < x) / \sum_{i=1}^{N} w_{x,i}$, where $w_{y,i} = 1 - \int_0^1 I((x_i, y) \in S)\, \mathrm{d}y$. In this figure, $w_{x,1}$ is the total length of the green part of the vertical line segments crossing over e_1, and $w_{x,2} = 1$ because the vertical line segment crossing e_2 has no intersection with S.

Because F_Y and F_X are unknown, we replace them with $\tilde{F}_Y$ and $\tilde{F}_X$, respectively; that is,

$$w_{x,i} = \frac{1}{1 - \int_0^1 I((x_i, y) \in S) \mathrm{d}\tilde{F}_y(y)}, \qquad w_{y,i} = \frac{1}{1 - \int_0^1 I((x, y_i) \in S) \mathrm{d}\tilde{F}_X(x)}. \tag{11}$$

The above equation, together with (7) and (8), form a solvable equation system. Below we propose an algorithm to solve this equation system.

First, the missing region S needs to satisfy the following condition.

Condition 1. *The projections of $([0, T] \times M) \setminus S$ (i.e., the sub-region in which no event is missing) on the t- and m-axes cover the entire observation period and the entire range of possible marks, respectively.*

This requirement ensures that the empirical distributions of $\{t_i\}$ and $\{m_i\}$ can be restored. With Condition 1 satisfied, when a record is incomplete, we can determine the area, say S, outside of which the record is complete. This

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

can be done either in the original time-mark plot, based on prior knowledge of the data quality, or in the BEPIT domain, based on the statistics R or D.

The algorithm to replenish the record includes three key steps: (1) transforming the process onto $[0,1]^2$ using the BEPIT to find a time–mark range that likely contains all missing events; (2) estimating a new empirical distribution function based on the data in the time–mark range, inside which events are supposed to be completely observed; (3) generating events in the missing region.

Initial settings. *Given the data set* $N_{\text{obs}} = \{(t_i, m_i) : i = 1, 2 \cdots, n\}$ *observed in* $[0, T] \times M$ *and a time–mark range* S*, known to include the missing events, suppose that* S *satisfies Condition 1.*

Step 1. We project the observed data and the range S *that contains the missing data onto* $[0,\, 1]^2$ *using the BEPIT in (4). Explicitly, set*

$$(t_i^{(1)}, m_i^{(1)}) = \Gamma_{N_{\text{obs}}}^{(1)}(t_i, m_i), \tag{12}$$

where

$$\Gamma_{N_{\text{obs}}}^{(1)}(t, m) = \left(\tilde{F}^{(1)}(t),\, \tilde{G}^{(1)}(m)\right) = \left(\frac{1}{n}\sum_{j=1}^{n} \mathbf{1}(t_j < t),\ \frac{1}{n}\sum_{j=1}^{n} \mathbf{1}(m_j < m)\right). \tag{13}$$

Denote $S^{(1)}$ *as the image of* S *under the transformation* $\Gamma_{N_{\text{obs}}}^{(1)}$.

Step 2. Starting from $\ell = 1$*, repeat the following iterative computation until convergence (e.g.,* $\max\{|t_i^{(\ell+1)} - t_i^{(\ell)}|, |m_i^{(\ell+1)} - m_i^{(\ell)}|\} < \epsilon$*), where* ϵ *is a given small positive number*:

$$(t_i^{(\ell+1)}, m_i^{(\ell+1)}) = \Gamma_{N_{\text{obs}}}^{(\ell+1)}(t_i^{(\ell)}, m_i^{(\ell)}; S^{(\ell)}), \quad i = 1, 2, \cdots, n, \tag{14}$$

$$S^{(\ell+1)} = \Gamma_{N_{\text{obs}}}^{(\ell+1)}(S^{(\ell)}; S^{(\ell)}), \tag{15}$$

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

where

$$\Gamma^{(\ell+1)}_{N_{\text{obs}}}(t,m;A)=\left(\frac{\sum_{j=1}^{n} w_1^{(\ell)}(t_j^{(\ell)},m_j^{(\ell)},A)\,\mathbf{1}(t_j^{(\ell)}<t)}{\sum_{j=1}^{n} w_1^{(\ell)}(t_j^{(\ell)},m_j^{(\ell)},A)},\frac{\sum_{j=1}^{n} w_2^{(\ell)}(t_j^{(\ell)},m_j^{(\ell)},A)\,\mathbf{1}(m_j^{(\ell)}<m)}{\sum_{j}^{n} w_2^{(\ell)}(t_j^{(\ell)},m_j^{(\ell)},A)}\right) \tag{16}$$

with the weights defined as

$$w_1^{(\ell)}(t,m,A)=\frac{\mathbf{1}\left((t,m)\notin A\right)}{1-\int_0^1 \mathbf{1}\left((t,m')\in A\right)\,\mathrm{d}G^{(\ell)}(m')} \tag{17}$$

$$w_2^{(\ell)}(t,m,A)=\frac{\mathbf{1}\left((t,m)\notin A\right)}{1-\int_0^1 \mathbf{1}\left((t',m)\in A\right)\,\mathrm{d}F^{(\ell)}(t')}, \tag{18}$$

for any regular region $A\subset[0,1]^2$. Denote the results upon convergence as $N^*_{\text{obs}}=\{(t_i^*,m_i^*):i=1,\,2\cdots,n\}$ and S^*.

Step 3. Generate a random number K from a negative binomial distribution, with parameters $(k,1-|S^*|)$, where $|S^*|$ is the area of S^* and

$$k=\sum_{i=1}^{n}\mathbf{1}((t_i^*,m_i^*)\notin S^*)=\#(N^*_{\text{obs}}\setminus S^*).$$

Step 4. Generate K random events independently, identically, and uniformly distributed in S^*. Denote these newly generated events as N^*_{rep}.

Step 5. For each event in N^*_{obs}, say, (t_j,m_j), that falls in S^*, sequentially remove from N^*_{rep} the event that is the closest to (t_j,m_j).

Step 6. Convert the resulting N^*_{rep} from the last step to the original observation space $[0,T]\times M$ through linear interpolation:

$$s_j=\mathrm{LI}\left(s_j^*;\,[0,\,t_1^*,\,t_2^*,\,\cdots,\,t_n^*,1],\,[0,\,t_1,\,t_2,\,\cdots,\,T]\right), \tag{19}$$

$$v_j=\mathrm{LI}\left(v_j^*;\,[0,\,m_1^*,\,m_2^*,\,\cdots,\,m_n^*],\,[0,\,m_1,\,m_2,\,\cdots,\,m_n]\right), \tag{20}$$

for each $(s_j^*,v_j^*)\in N^*_{\text{rep}}$, where $\mathrm{LI}(x,A,B)$ represents the linear interpolation value of x, conditional on the function values for each component

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

in A being locations corresponding to each component in B. Denote the set consisting of all (s_j, v_j) as N_{rep}.

Final output. *Return N_{rep}.*

Example 4. *Here we apply the above algorithm to the thinned data set in Example 2. The output from Steps 4 to 6 is shown in Figures 4(b)–(c). The final output for our simulation example is shown in Figure 4(d). The tests using statistics R and D in (6) give p-values of 0.605 and 0.718, respectively, providing no evidence to reject the hypothesis that the replenished data set is complete (Figure 4(e)). Figure 4(f) compares the cumulative numbers of events in the original, observed, and replenished processes, showing that the replenishing algorithm recovers the missing data to some extent.*

Notes:

(1) Equation (13) is the BEPIT mentioned in the previous section. If the data are completely recorded, $\{(t_i^{(1)}, m_i^{(1)}), i = 1, 2 \cdots, n\}$ form an approximately homogeneous process on $[0, 1]^2$. As shown in Figure 2(b), the sparseness of the points around the lower, left corner implies that smaller events are missing in the earlier period. Rather than choosing S in Figure 1(a), it is more convenient to specify $S^{(1)}$ directly in Figure 2(a) or (b).

(2) Step 2 is carried out based on the fact that the transformation $\Gamma_{N_{\text{obs}}}$ and $S^{(1)} = \Gamma_{N_{\text{obs}}}(S)$ can be quite different from Γ_N, owing to the missing data. The iteration in this step helps us construct a bi-scale transformation as close as possible to the BEPIT yielded by the complete data

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

(i.e., $\Gamma^*_{N_{\text{obs}}} \approx \Gamma_N$). At the same time, the corresponding area that contains the missing data, S^*, is restored. This can be seen by comparing Figures 1(b) and 4(b).

Step 2 essentially solves F^* and G^* in the following equations:

$$F^*(t) = \frac{\sum_{j=1}^n w_1(t_j, m_j, S)\, \mathbf{1}\,(t_j < t)}{\sum_{j=1}^n w_1(t_j, m_j, S)}, \tag{21}$$

$$G^*(m) = \frac{\sum_{j=1}^n w_2(t_j, m_j, S)\, \mathbf{1}\,(m_j < m)}{\sum_{j=1}^n w_2(t_j, m_j, S)}, \tag{22}$$

where

$$w_1(t, m, S) = \frac{\mathbf{1}\,((t, m) \notin S)}{1 - \int_M \mathbf{1}\,((t, m') \in S)\, \mathrm{d}G^*(m')} \tag{23}$$

$$w_2(t, m, S) = \frac{\mathbf{1}\,((t, m) \notin S)}{1 - \int_M \mathbf{1}\,((t', m) \in S)\, \mathrm{d}F^*(t')}. \tag{24}$$

If we define $\Gamma^*_{N_{\text{obs}}}(t, m) = (F^*(t), G^*(m))$ as a mapping from $[0, T] \times M$ to $[0, 1]^2$, then $\Gamma^*_{N_{\text{obs}}}(t, m)$ directly maps N_{obs} to N^*_{obs} and S to S^*.

The existence of a solution in the iteration given by (21) to (24) and the asymptotic property of the solution are given in the Supplementary Material.

(3) Steps 3 and 4 are based on the following fact: given a homogeneous Poisson process with an unknown occurrence rate, if there are k events falling within an area of S_1, then the number of events falling in the complementary area, S_2, follows a negative binomial distribution with parameter $(k, |S_1|/(|S_1| + |S_2|))$ (e.g., DeGroot, 1986, 258–259).

(4) In Step 5, we should keep the existing events observed in S, and remove the same number of simulated points.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

One advantage of the algorithm is that if S is unknown, we can use the time-mark plot of $N^{(1)}$, as in Figure 2(b), to determine $S^{(1)}$ by justifying which region is likely to contain the missing events, and then continue with Step 2. Once the replenishment is complete, S can be obtained by substituting the coordinate of each point on the boundary of S^* into (19) and (20).

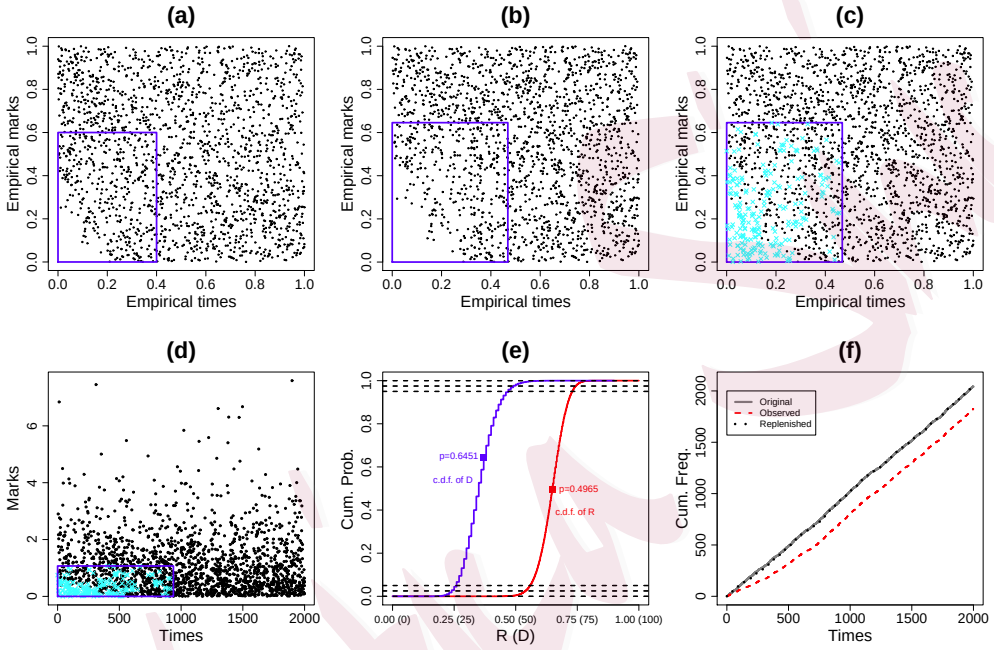

Figure 4: An application of the proposed replenishing algorithm to the synthetic data set. (a) Rescaled marks versus rescaled occurrence times of the observed events (dots), with the bi-scale transformation $\Gamma_{N_{\text{obs}}}$ based on the observed process. The polygon is the missing area, $S^{(1)}$. (b) Rescaled marks versus rescaled occurrence times of the observed events (dots), with the rescaling $\Gamma^*_{N_{\text{obs}}}$ based on the events outside of S. The polygon is the missing area after transformation $\Gamma^*_{N_{\text{obs}}}$, that is., S^*. (c) Rescaled marks versus rescaled occurrence times of the observed and replenished events (crosses) (i.e., newly generated events after removing events that are closest to any of those observed in S, with the rescaling $\Gamma^*_{N_{\text{obs}}}$ based on the empirical distributions of the events outside S. (d) Marks versus occurrence times of the observed synthetic events and the replenished events. (e) Cdfs of R and D for testing missing data in the replenished data set in (c). (f) Cumulative frequencies versus occurrence times for the original, observed, and replenished processes.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

2.4 Additional simulations

To illustrate the overall behavior of the above replenishing algorithm, we repeat the algorithm many times, with S fixed, for the following two cases: (1) simulating a Poisson process with $\lambda = 2,000$; and (2) simulating Poisson processes with rate λ, drawn from a uniform distribution within [100, 3,000]. Both simulations have the same missing probability functions, as given by (5). Figures 5(a) and (b) compare between the true numbers of missing events and the numbers of replenished events for cases (1) and (2), respectively. In Figure 5(a), since λ is fixed, the number of replenished events is independent of the true number of missing events, and has a larger variance. Several statistics related to these simulations are given in Table 1, including the mean numbers and the variances of the missing and replenished points, the mean of the relative differences, and the relative difference between the means in 500 and 2,000 simulations. In particular, the near-zero relative deviation of the mean number of replenished events shows that the proposed method is consistent. Here, the larger values of the mean relative deviation of the number of replenished events from the number of missing events illustrate the nature of the uncertainty related to the problem. Such uncertainty is produced not only by the randomness of the numbers of replenished and missing events, but also by the uncertainty in the estimation of the occurrence rate in the process from the events in the nonmissing part. In Figure 5(b), the expected number of replenished events in many repeated simulations is close to the number of missing events. Moreover, the relative deviation decreases when the number of missing events (or λ) increases. These results imply that this algorithm replenishes the missing events reasonably well. In addition, when λ or the number of events in the process is quite small, some

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Table 1: Statistics related to Figure 5(a). $\#m$: number of missing points; $\#r$: number of replenished points; $\bar{\cdot}$: mean value; $\sigma(\cdot)$: standard deviation.

#simu.	$\overline{\#m}$	$\sigma(\#m)$	$\overline{\#r}$	$\sigma(\#r)$	$\overline{\left[\frac{\lvert\#m-\#r\rvert}{\#m}\right]}$	$\frac{\lvert\overline{\#m}-\overline{\#r}\rvert}{\overline{\#m}}$
500	228.274	14.929	232.006	63.926	0.226	0.016
2000	227.712	14.719	230.860	62.145	0.224	0.014

outputs yield a negative number of replenished events (when the number of missing events is less than 50 in Figure 5(b)). The number of replenished events is calculated simply as the number of simulated events in S in Steps 3 and 4 minus the number of observed events in S. This finding indicates that the existence of missing data in these situations cannot be quantified probabilistically.

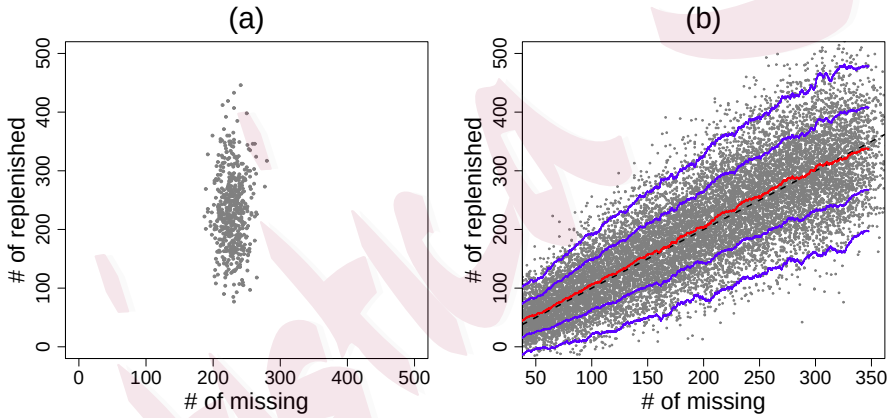

Figure 5: Comparison between the number of true missing events and the number of replenished events. (a) $\lambda = 2,000$, fixed. (b) λ is drawn from a uniform distribution between 100 and 3,000. The dashed line represents the case where the numbers of missing and replenished events are equal. The curves represent the running mean and the corresponding single and double standard deviation bands.

3 Application

3.1 Volcanic eruption record

In this example, we analyze a record of the eruptions from the Hakone volcano, an active volcano located at the northern boundary zone of the

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

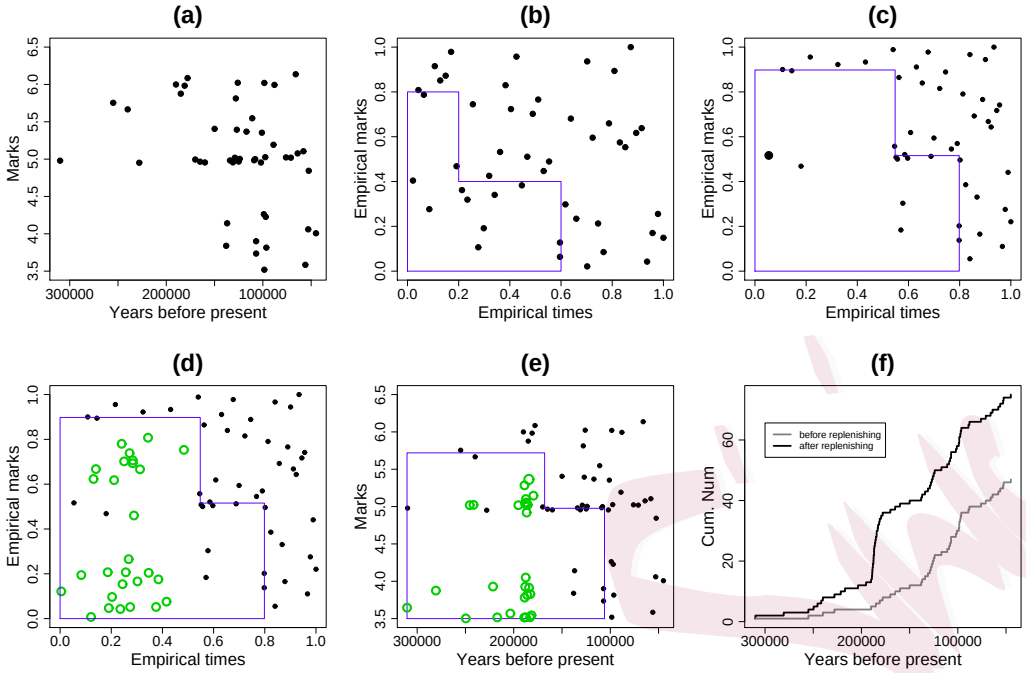

Figure 6: Results after applying the replenishment algorithm to volcanic eruption data. (a) Marks versus occurrence times of the eruption events. (b) Empirical distribution of marks versus that of occurrence times. (c) Rescaled marks versus rescaled occurrence times, with the rescaling based on the empirical distributions of the events outside of S. (d) Rescaled marks versus rescaled occurrence times of the observed and replenished events (i.e., newly generated events after removing events that are closest to any of those observed in S), with the rescaling based on the empirical distributions of the events outside of S. (e) Marks versus occurrence times of the observed and replenished events. (f) Cumulative numbers of events against occurrence times. The polygon is the area S and its corresponding mappings, in which the missing events fall. The green dots are the replenished events.

Izu-Mariana volcanic arc in central Japan (Yukutake et al., 2010; Honda et al., 2014). The data on Japanese explosive eruptions are compiled from the Smithsonian's Global Volcanism Program database (Siebert and Simkin, 2002), the Large Magnitude Explosive Volcanic Eruptions database (LaMEVE database, Crosweller et al., 2012), and additional Japanese databases (Machida and Arai, 2003; Committee for Catalog of Quaternary Volcanoes in Japan (ed), 2000; Geological Survey of Japan, AIST (ed), 2013; Hayakawa, 2010).

For the Hakone volcano, 46 of the 54 compiled events have an eruption

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

magnitude ($M = \log_{10}$[erupted mass in kg] $- 7$; see Pyle (2015)) equal to or larger than 4.0 (Table S1 in Supplementary material). Figure 6(a) shows the eruption magnitudes versus occurrence times of these 46 events. Figure 6(b) shows the empirical distribution, transformed following Step 1 of the algorithm. From this plot, the polygon boundaries of S are determined based on the following assumptions. First, the events of empirical marks < 0.8 ($M < 5.7$) are missing before the empirical time $= 0.2$ (165 ka). Second, the recording of larger events improves after the empirical time $=$ 0.2 (165 ka), although the events of empirical marks < 0.4 ($M < 5.0$) are still missing. Third, the recording of events improves further and there are no missing events after the empirical time $= 0.6$ (105 ka). The results of the replenishing algorithm are shown in Figures 6(c) to 6(e).

The estimated cumulative number of events for the replenished data set shows a remarkable jump of around 180 ka (Figure 6(f)). This jump is caused by the replenished events synthesized around 180 ka (Figure 6(e)), based on the cluster of four large events ($M \sim 6$) at 178 ka, 181 ka, 185 ka, and 190 ka (Figure 6(a); Hayakawa, 2010). The ages of the events at the Hakone volcano are still not fully agreed upon in the literature. For example, Yamamoto (2015) assumed that the ages of the aforementioned eruptions are about 135 ka, 135 ka, 180 ka, and 215 ka, respectively. Therefore, the reliability of the jump of the cumulative number of events (Figure 6(f)) might problematic in the volcanological dating of event ages, as in estimating the tephra volume and rounded eruption magnitude in volcanology (Brown et al., 2014). For example, the analyzed data set has clusters of events of magnitudes 4 and 5 (Figure 6(a)) and, therefore, the replenished events around 180 ka are also clustered around magnitudes 4 and 5 (Figure 6(e)).

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Note that it is difficult to determine the exact period of under-recording in the eruption history of each volcano. Kiyosugi et al. (2015) showed that many eruptions are still missing from the overall Japanese database, even for the last 100,000 years. Therefore, the polygon shape (Figure 6(b)) we have used suggests that our replenished data have the same completeness level as that of the data outside the polygon. Our method is a way to consider the under-recording of events in volcanic hazard assessments of explosive eruptions using geological records.

3.2 Earthquake catalog: Missing aftershocks

It is well known that immediately after a large earthquake, many aftershocks cannot be recorded, because the seismic waveforms generated by the aftershocks, many of which occur in a short time after the mainshock, overlap with each other and cannot be distinguished. In this section, we study the earthquake catalog from Southwest China for the period from January 1, 1990 to April 20, 2013, in a space range of $26^\circ - 34^\circ N$ and $97^\circ - 107^\circ E$, with minimum magnitude 3.0 (Figure S2 in the Supplementary Material). This data set is selected from the Chinese Earthquake catalog compiled by the China Earthquake Data Center (CEDC) (URL: http://data.earthquake.cn/index.html). The Wenchuan Mw 7.9 (Ms 8.0) earthquake, which occurred on May 12, 2008, was one of the two largest seismic events in China during the last 50 years. There are 6,249 events in the selected space and time range, of which 3,754 occurred after the Wenchuan earthquake, indicating a low seismicity level above magnitude 3.0 in the study region prior to 2008. Many aftershocks are missing immediately after the mainshock. In particular, events of magnitudes between 3.0 and 4.0 are not properly recorded for a period

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

of about one-and-a-half months after the mainshock. The majority of the events after May 12, 2008, can be taken as clustering events triggered by the Wenchuan mainshock. When analyzing the seismicity in this area, Jia et al. (2014) and Guo et al. (2015) chose a relatively high magnitude threshold of 4.0 to avoid biases in estimates caused by missing events. As a results, 5,217 of the 6,249 events had to be ignored.

This example is quite different from the previous example and that based on the simulated data. The missing range can be well specified before replenishment: the missing values are known immediately after the occurrence of the mainshock, and the monitoring ability for events between magnitudes 3.0 and 4.0 are restored one-and-a-half months later. The results are illustrated in Figure 7. We can see that missing events take up about half the total number of events.

In seismology, the frequency of aftershock occurrences in an aftershock sequence can be modeled by the empirical Omori–Utsu formula (e.g., Utsu et al., 1995),

$$\lambda(t) = \frac{K}{(t+c)^p}, \tag{25}$$

where K is an index proportional to the number of earthquakes excited by the mainshock, c is related to the period after the mainshock from which the aftershock rate drops slowly, and p is the power related to the decay rate of the aftershocks. Utsu et al. (1995) discussed how the parameters c and p change with the cut-off magnitude threshold, and hypothesized that such changes occur because small aftershocks in an early stage of the sequence are missing from the catalog. We fit the above Omori–Utsu formula to both the original and the replenished catalogs (Table 2), and obtain maximum

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

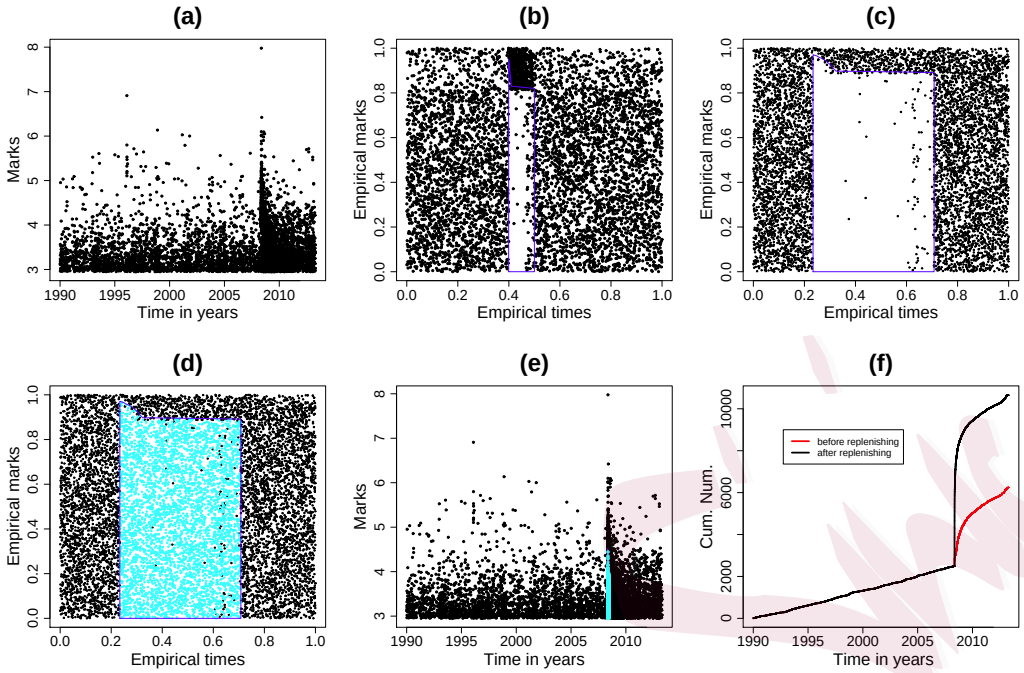

Figure 7: Results from applying the replenishment algorithm to earthquake data from Southwest China. (a) Marks versus occurrence times of the earthquake events. (b) Empirical distribution of marks versus that of occurrence times. (c) Rescaled marks versus rescaled occurrence times, with the rescaling based on the empirical distributions of the events outside of S. (d) Rescaled marks versus rescaled occurrence times of the observed and replenished events (i.e., newly generated events after removing events that are closest to any of the observed in S), with the rescaling based on the empirical distributions of the events outside S. (e) Marks versus occurrence times of the observed and replenished events. (f) Cumulative numbers of events against occurrence times. The polygon is the area S and its corresponding mappings, in which the missing events fall.

likelihood estimates of the parameters. The results show that after the replenishment, the Omori parameters c and p no longer change. We also fit the Omori–Utsu formula to the original data set, but only consider earthquakes that occurred at least 54 days after the mainshock. In this case, although c and p are slightly different from the estimates for the replenshed data from the starting time, they do not change much when the magnitude threshold changes from 2.95 to 4.15 (Table S2 in the Supplementary Material). These results numerically confirm the hypothesis of Utsu et al. (1995) that missing

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

Magnitude threshold	Replenished dataset $[t_{\text{main}}, T]$			Orig. dataset $[t_{\text{main}}, T]$		
	$\hat{K}$	$\hat{c}$	$\hat{p}$	$\hat{K}$	$\hat{c}$	$\hat{p}$
2.95	804.4	.1140	1.003	82.29	.0553	.6205
3.05	639.2	.1131	1.003	80.31	.0596	.6547
3.15	511.5	.1134	1.001	79.25	.0660	.6872
3.25	412.9	.1110	.9965	79.04	.0737	.7185
3.35	327.3	.1067	.9926	78.80	.0825	.7555
3.45	260.3	.1141	.9925	80.67	.0991	.7986
3.55	213.8	.1142	.9953	83.33	.1177	.8407
3.65	171.6	.1135	.9907	85.73	.1360	.8799
3.75	135.9	.1132	.9911	90.18	.1642	.9278
3.85	111.2	.1029	.9941	95.17	.1935	.9708
3.95	100.0	.1241	1.015	103.2	.2383	1.023
4.05	74.12	.1082	1.013	79.20	.1938	1.027
4.15	60.65	.1266	1.026	62.92	.1690	1.034

Table 2: Results from fitting the Omori–Utsu formula to the original and the replenished data sets of earthquakes from Southwest China, with different magnitude thresholds. t_{main}: occurrence time of the mainshock; T: end of the time interval.

small events in the early stage of an aftershock sequence causes the instability of the estimate of the Omori–Utsu formula.

4 Conclusion

In this study, we proposed a method for replenishing missing data in marked temporal point processes, based only on the assumption that the marks of the events are separable from the occurrence times, regardless of how the events interact on the time axis. The key point of this method is an algorithm that iteratively estimates the missing area in the transformed domain, based on the parts where the data are completely recorded. We applied the proposed method to an eruption record of the Hakone volcano in Japan and to an earthquake catalog from Southwest China, which includes the aftershock zone of the 2008 Mw7.9 Wenchuan earthquake. The results show that the

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

proposed method helps to both evaluate the influence of missing data and correct the bias caused by such data.

Detection of the missing area In our two examples, the missing area is determined by visual inspection of the bi-scale transformed data for the historical records of the Hakone volcano, and by prior information on the seismic network for the Wenchuan aftershock sequence. In most cases, the missing area is determined by the experience of data analysts or by information on the data from other sources. However, it is possible to turn the replenishing algorithm into an automated algorithm.

Starting from $S' = \emptyset$, we divide the unit square into small cells in the bi-scale transformed domain, obtained by applying the transformation defined in (9) to (13). Then, we carry out the statistical tests based on the statistics R or D on the cells that do not intersect S', as discussed in Section 3. If the test shows that missing cells exist, then we merge these cells into S'. These steps are iterated until no further cells are added to S'. Note that because this topic belongs within the scope of data processing algorithms, we did not include it in this paper.

Separability of marks As discussed earlier, the applicability of this algorithm depends on whether the mark distribution is separable from the occurrence time. If such dependence is known explicitly as a probability density function, say $g(m \,|\, t)$, we can directly use the cdf that corresponds to f in Steps 1 and 2 in the algorithm (i.e., $m_i^{(\ell)} = G(m_i \,|\, t_i)$, for $\ell \geq 1$). Of course, such dependence should also be considered when transforming the marks of replenished events from $[0, 1]$ to the original mark space. If

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

the mark is dependent on time, but we do not know how, together with the existence of missing events, the replenishment/imputation problem becomes unidentifiable.

Another case worth discussing is when the mark distribution is known and does not depend on time. We can again use the cdf of the marks in Steps 1 and 2 directly in the algorithm (i.e., by setting $m_i^{(\ell)} = G(m_i)$, for $\ell \geq 1$). Such missing data can also be estimated using Bayesian methods, as in Ogata and Katsura (1993), and then replenished by direct simulation.

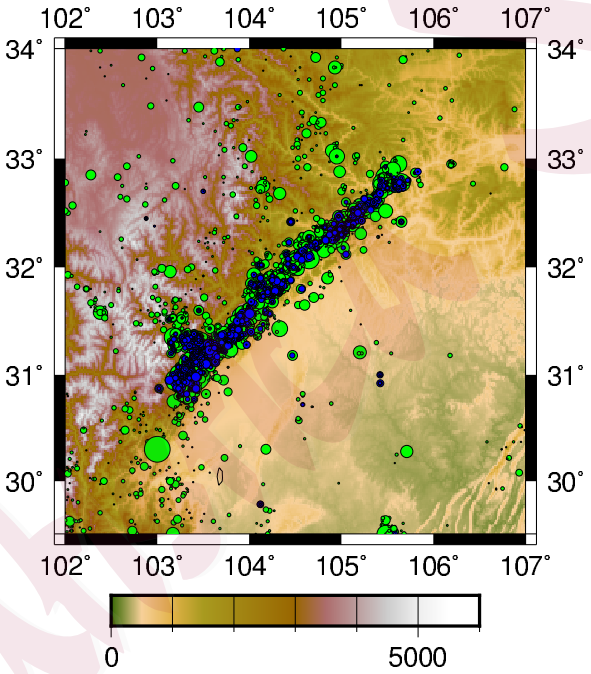

Figure 8: Epicenter map of imputed earthquakes (solid blue circles) for the Wenchuan aftershock sequence.

Imputation of locations This method is powerful for marked temporal point processes, but it cannot be extended easily to high-dimensional or spatiotemporal cases because, in most cases, the process is not homogeneous in space. However, it is still possible case by case. For example, to replenish the Wenchuan aftershock sequence, we can use the clustering feature of earth-

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

quakes. A simple replenishing algorithm is as follows. For each simulated event, find a fixed number (e.g., 50) of events closest to it in time in the observed process. Then, construct a Delaunay tessellation network for these 50 events, and select with equal probability one of the Delaunay triangles. Lastly, place the the simulated event randomly and uniformly in this selected triangle. An example of the imputed locations of the missing aftershocks of the Wenchuan earthquake is shown in Figure 8. For a spatially inhibitive process, different methods should be used.

In summary, the proposed method is useful when dealing with the missing data problem in point-process observations, such as volcano eruption records and historical or short-term earthquake catalogs.

Supplementary Material

The online Supplementary Material includes the following topics: (1) a proof of the existence of a solution to the equation system in (21) to (24); (2) the asymptotic properties of the solution; (3) additional simulations for the case in which the missing region is wrongly specified; (4) list of the history record of the Hakone volcano; and (5) comments on the Wenchuan aftershock sequence.

Acknowledgements

This project was supported by the Royal Society of New Zealand Marsden Fund (contact UOO1419). JZ was also partially supported by Grants-in-Aid No. 2530052 for Scientific Research (C) from the Japan Society for the Promotion of Science. Helpful discussions with David Harte from GNS Science

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

New Zealand and Boris Baeumer from the University of Otago are gratefully acknowledged. The authors also thank the AE and three anonymous reviewers for their encouragement and constructive comments.

References

Bebbington, M. S. (2014). Long-term forecasting of volcanic explosivity. *Geophysical Journal International*, 197:1500–1515.

Brown, S. K., Crosweller, H. S., Sparks, R. S. J., Cottrell, E., Deligne, N. I., Guerrero, N. O., Hobbs, L., Kiyosugi, K., Loughlin, S. C., Siebert, L., and Takarada, S. (2014). Characterisation of the quaternary eruption record: analysis of the large magnitude explosive volcanic eruptions (LaMEVE) database. *J Appl Volcanol*, 3:5.

Committee for Catalog of Quaternary Volcanoes in Japan (ed) (2000). Catalog of quaternary volcanoes in Japan (in Japanese).

Corrado, C. J. (2011). The exact distribution of the maximum, minimum and the range of multinomial/dirichlet and multivariate hypergeometric frequencies. *Statistics and Computing*, 21(3):349–359.

Crosweller, H., Arora, B., Brown, S. K., Cottrell, E., Deligne, N., Guerrero, N., Hobbs, L., Kiyosugi, K., C., L. S., Lowndes, J., Nayembil, M., Siebert, L., Sparks, R. S. J., Takarada, S., and Venzke, E. (2012). Global database on large magnitude explosive volcanic eruptions (LaMEVE). *Journal of Applied Volcanology*, 1:4.

Daley, D. D. and Vere-Jones, D. (2003). *An Introduction to Theory of Point Processes – Volume 1: Elementary Theory and Methods (2nd Edition)*. Springer, New York, NY.

Daley, D. D. and Vere-Jones, D. (2008). *An Introduction to Theory of Point Processes – Volume II: General Theory and Structure (2nd Edition)*. Springer, New York, NY.

DeGroot, M. H. (1986). *Probability and Statistics (Second ed.)*. Addison-Wesley.

Diggle, P. J. and Rowlingson, B. S. (1994). A conditional approach to point process modelling of elevated risk. *Journal of the Royal Statistical Society. Series A (Statistics in Society)*, 157(3):433–440.

Enescu, B., Mori, J., and Miyazawa, M. (2007). Quantifying early aftershock activity of the 2004 mid-Niigata prefecture earthquake ($Mw6.6$). *Journal of Geophysical Research: Solid Earth*, 112(B4):B004629.

Enescu, B., Mori, J., Miyazawa, M., and Kano, Y. (2009). Omori-Utsu law c-values associated with recent moderate earthquakes in Japan. *Bulletin of the Seismological Society of America*, 99(2A):884–891.

Geological Survey of Japan, AIST (ed) (2013). Catalog of eruptive events during the last 10,000 years in japan, version 2.1 (in japanese). Technical report.

Guo, Y., Zhuang, J., and Zhou, S. (2015). An improved space-time ETAS model for inverting the rupture geometry from seismicity triggering. *Journal of Geophysical Research: Solid Earth*, 120(5):3309–3323. 2015JB011979.

Gutenberg, B. and Richter, C. F. (1944). Frequency of earthquakes in California. *Bull. Seis. Soc. Am.*, 34:184–188.

578 Hainzl, S. (2016). Rate-dependent incompleteness of earthquake catalogs.
579 *Seismological Research Letters*, 87(2A):337–344.

580 Hayakawa, Y. (2010). Hayakawafs 2000-year eruption database and one
581 million-year tephra database, http://www.hayakawayukio.jp/database/.

582 Honda, R., Yukutake, Y., Yoshida, A., Harada, M., Miyaoka, K., and Sato-
583 mura, M. (2014). Stress-induced spatiotemporal variations in anisotropic
584 structures beneath hakone volcano, japan, detected by s wave splitting:
585 A tool for volcanic activity monitoring. *Journal of Geophysical Research:*
586 *Solid Earth*, 119(9):7043–7057.

587 Iwata, T. (2008). Low detection capability of global earthquakes after the
588 occurrence of large earthquakes: Investigation of the Harvard CMT cata-
589 logue. *Geophysical Journal International*, 174(3):849–856.

590 Iwata, T. (2013). Estimation of completeness magnitude considering daily
591 variation in earthquake detection capability. *Geophysical Journal Interna-*
592 *tional*, 194(3):1909–1919.

593 Iwata, T. (2014). Decomposition of seasonality and long-term trend in seis-
594 mological data: A Bayesian modelling of earthquake detection capability.
595 *Australian & New Zealand Journal of Statistics*, 56(3):201–215.

596 Jia, K., Zhou, S., Zhuang, J., and Jiang, C. (2014). Possibility of the inde-
597 pendence between the 2013 Lushan earthquake and the 2008 Wenchuan
598 earthquake on Longmen Shan Fault, Sichuan, China. *Seismological Re-*
599 *search Letters*, 85(1):60–67.

600 Johnson, N. L. (1960). An approximation to the multinomial distribution
601 some properties and applications. *Biometrika*, 47(1-2):93–102.

Johnson, N. L. and Young, D. H. (1960). Some applications of two approximations to the multinomial distribution. *Biometrika*, pages 463–469.

Karr, A. (1991). *Point Processes and Their Statistical Inference*. Marcel Dekker, Inc., New York and Basel.

Kiyosugi, K., Connor, C. B., Sparks, R. S. J., Crosweller, H. S., Brown, S. K., Siebert, L., Wang, T., and Takarada, S. (2015). How many explosive eruptions are missing from the geologic record? analysis of the quaternary record of large magnitude explosive eruptions in japan. *Journal of Applied Volcanology*, 4:17.

Machida, H. and Arai, F. (2003). *Atlas of Tephra in and around Japan, revised edition*. University of Tokyo Press, Japan (in Japanese).

Marsan, D. and Enescu, B. (2012). Modeling the foreshock sequence prior to the 2011, MW9.0 Tohoku, Japan, earthquake. *Journal of Geophysical Research: Solid Earth*, 117(B6):B06316.

Mohler, G. O., Short, M. B., Brantingham, P. J., Schoenberg, F. P., and Tita, G. E. (2011). Self-exciting point process modeling of crime. *Journal of the American Statistical Association*, 106(493):100–108.

Møller, J. and Waagepetersen, R. P. (2003). *Statistical Inference and Simulation for Spatial Point Processes*. Chapman and Hall/CRC.

Ogata, Y. (1988). Statistical models for earthquake occurrences and residual analysis for point processes. *Journal of the American Statistical Association*, 83(401):9–27.

Ogata, Y. (2006). Monitoring of anomaly in the aftershock sequence of the 2005 earthquake of M7.0 off coast of the western Fukuoka, Japan, by the ETAS model. *Geophysical Research Letters*, 33:L01303.

Ogata, Y. and Katsura, K. (1993). Analysis of temporal and spatial heterogeneity of magnitude frequency distribution inferred from earthquake catalogues. *Geophysical Journal International*, 113(3):727–738.

Ogata, Y., Katsura, K., Falcone, G., Nanjo, K., and Zhuang, J. (2013). Comprehensive and topical evaluations of earthquake forecasts in terms of number, time, space, and magnitude. *Bulletin of the Seismological Society of America*, 103(3):1692–1708.

Ogata, Y. and Vere-Jones, D. (2003). Examples of statistical models and methods applied to seismology and related earth physics. In Lee, W. H., Kanamori, H., Jennings, P. C., and Kisslinger, C., editors, *International Handbook of Earthquake and Engineering Seismology, Vol.81B*, chapter 82. International Association of Seismology and Physics of Earth's Interior.

Ogata, Y. and Zhuang, J. (2006). Space-time ETAS models and an improved extension. *Tectonophysics*, 413(1-2):13–23.

Omi, T., Ogata, Y., Hirata, Y., and Aihara, K. (2013). Forecasting large aftershocks within one day after the main shock. *Scientific Reports*, 3:2218.

Omi, T., Ogata, Y., Hirata, Y., and Aihara, K. (2014). Estimating the ETAS model from an early aftershock sequence. *Geophysical Research Letters*, 41(3):850–857.

Omi, T., Ogata, Y., Hirata, Y., and Aihara, K. (2015). Intermediate-term forecasting of aftershocks from an early aftershock sequence: Bayesian and ensemble forecasting approaches. *Journal of Geophysical Research: Solid Earth*, 120(4):2561–2578.

Passarelli, L., Sandri, L., Bonazzi, A., and Marzocchi, W. (2010). Bayesian hierarchical time predictable model for eruption occurrence: an application to kilauea volcano. *Geophysical Journal International*, 181(3):1525–1538.

Peng, Z., Vidale, J. E., Ishii, M., and Helmstetter, A. (2007). Seismicity rate immediately before and after main shock rupture from high-frequency waveforms in Japan. *Journal of Geophysical Research: Solid Earth*, 112(B3):B03306.

Pyle, D. M. (2015). Sizes of volcanic eruptions. In Sigurdsson, H., editor, *The Encyclopedia of Volcanoes (Second Edition)*, chapter 13, pages 257 – 264. Academic Press, Amsterdam, second edition edition.

Sawazaki, K. and Enescu, B. (2014). Imaging the high-frequency energy radiation process of a main shock and its early aftershock sequence: The case of the 2008 Iwate-Miyagi Nairiku earthquake, Japan. *Journal of Geophysical Research: Solid Earth*, 119(6):4729–4746.

Schoenberg, F. P. (2003). Multidimensional residual analysis of point process models for earthquake occurrences. *J. Am. Stat. Assoc.*, 98:789–795(7).

Schoenberg, F. P., Chang, C., Keeley, J., Pompa, J., Woods, J., and H., X. (2007). A critical assessment of the burning index in Los Angeles County, California. *International Journal of Wildland Fire*, 16:473?483.

Statistica Sinica: Preprint
doi:10.5705/ss.202017.0403

669 Siebert, L. and Simkin, T. (2002). Volcanoes of the world: an illustrated
670 catalog of holocene volcanoes and their eruptions, smithsonian institution,
671 global volcanism program digital information series, gvp-3.

672 Utsu, T., Ogata, Y., and Matsu'ura, R. S. (1995). The centenary of the
673 Omori formula for a decay law of aftershock activity. *Journal of Physics*
674 *of the Earth*, 43(1):1–33.

675 Vere-Jones, D. (1970). Stochastic models for earthquake occurrence. *J. Roy.*
676 *Stat. Soc. Series B (Methodological)*, 32(1):1–62 (with discussion).

677 Wang, T. and Bebbington, M. (2012). Estimating the likelihood of an erup-
678 tion from a volcano with missing onsets in its record. *Journal of Volcanol-*
679 *ogy and Geothermal Research*, 243–244:14–23.

680 Wang, T. and Bebbington, M. (2013). Robust estimation for the weibull
681 process applied to eruption records. *Mathematical Geosciences*, 45(7):851–
682 872.

683 Werner, M. J., Helmstetter, A., Jackson, D. D., and Kagan, Y. Y. (2011).
684 High-resolution long-term and short-term earthquake forecasts for Califor-
685 nia. *Bulletin of the Seismological Society of America*, 101(4):1630–1648.

686 Yamamoto, T. (2015). Cumulative volume step-diagrams for eruptive mag-
687 mas from major quaternary volcanoes in japan. Technical Report GSJ
688 Open-File Report, No.613, Geological Survey of Japan, AIST.

689 Yukutake, Y., Tanada, T., Honda, R., Harada, M., Ito, H., and Yoshida,
690 A. (2010). Fine fracture structures in the geothermal region of Hakone
691 volcano, revealed by well-resolved earthquake hypocenters and focal mech-
692 anisms. *Tectonophysics*, 489:104–118.

693 Zhuang, J. (2011). Next-day earthquake forecasts by using the ETAS model.
694 *Earth, Planet, and Space*, 63:207–216.

695 Zhuang, J., Ogata, Y., and Vere-Jones, D. (2002). Stochastic declustering
696 of space-time earthquake occurrences. *Journal of the American Statistical*
697 *Association*, 97(3):369–380.

698 Zhuang, J., Ogata, Y., and Vere-Jones, D. (2004). Analyzing earthquake clus-
699 tering features by using stochastic reconstruction. *Journal of Geophysical*
700 *Research*, 109(3):B05301.

701 Zhuang, J., Ogata, Y., and Wang, T. (2017). Data completeness of the Ku-
702 mamoto earthquake sequence in the JMA catalog and its influence on the
703 estimation of the ETAS parameters. *Earth, Planets and Space*, 69(1):36.

704 Zipkin, J. R., Schoenberg, F. P., Coronges, K., and Bertozzi, A. L. (2015).
705 Point-process models of social network interactions: Parameter estimation
706 and missing data recovery. *European Journal of Applied Mathematics*,
707 FirstView:1–28.